# ON THE DYNAMICS OF A DRIFT FLOW UNDER LOW WIND

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# **OBJECTIVE**

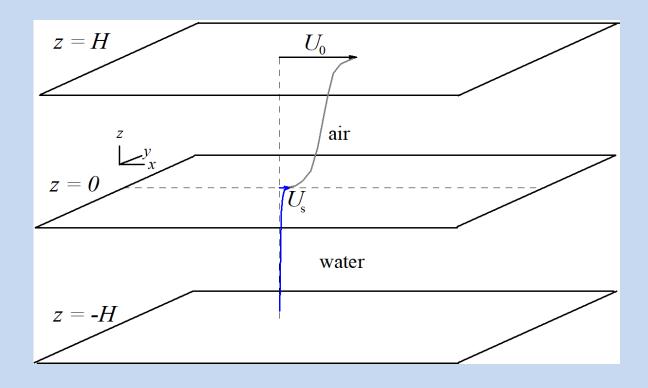
The studies (laboratory, numerical) of the wind-induced surface drift flow have been concerned mostly with moderate to strong winds (form 5 to 10 m/s) (*Banner & Peirson* 1998; *Cheung & Street R.L* 1988; *Longo et al.* 2012; *Polnikov et al.* 2020). This regime is characterized by the presence of surface (breaking and/or microbreaking) waves and underwater turbulence, contributing to the drift. The well-known parameterization of the surface drift under this "wave+turbulence" regime is (*Wu* 1975)

#### $U_s \approx 0.53 u_*$

A few studies (*Gemmrich & Hasse* 1992; *Shrira et al.* 2005) considered a sufficiently low-wind condition (below 2-3 m/s) where the water surface remains smooth. However, the question still stands how to parameterize the surface drift if the water flow remains quasi-laminar and the surface remains aerodynamically smooth.

This study addresses this problem by performing Direct Numerical Simulation (DNS)

### **Scheme of numerical experiment**



Cartesian coordinates are considered, with coordinates x –along the mean wind direction (bulk velocity U<sub>0</sub>, y – spanwise and z – vertical directions;  $0 \le x/H \le 6$ ,  $0 \le y/H \le 4$ ,  $-1 \le z/H \le 1$  discretized on staggered, z-nonuniform grid of  $360 \times 240 \times 360$  nodes. DNS is performed simultaneously in two layers:  $-H \le z \le 0$  – water, and  $0 \le z \le H$  – air with an interface (water surface) at z = 0. Dirichlet (no-slip) conditions are put at the bottom (at rest, z = -H) and top (moving with U<sub>0</sub>, z = H) boundary planes. All fields are periodical in x and y. At the water surface (z = 0) the continuity of momentum flux and mass conditions are imposed. The water surface is assumed to be <u>flat</u> throughout simulations.

## **Mathematical model**

The Navier-Stokes equations are solved for air and water:

$$\frac{\partial U_i^{a,w}}{\partial t} + \frac{\partial}{\partial x_j} \left( U_i^{a,w} U_j^{a,w} \right) = -\frac{\partial P_{a,w}}{\partial x_i} + \frac{1}{\operatorname{Re}_{a,w}} \frac{\partial^2 U_i^{a,w}}{\partial x_j^2}, \qquad (1)$$

$$\frac{\partial U_j^{a,w}}{\partial x_i} = 0. \qquad (2)$$

 $U_i^{a,w}$  and  $P_{a,w}$  - velocity and pressure in air (*a*) and water (*w*) (*i* = *x*,*y*,*z*),

$$\operatorname{Re}_{a,w} = \frac{U_0 H}{V_{a,w}} , \qquad (3)$$

is the Reynolds number,  $v_{a,w}$  - kinematic viscosities,  $8 \cdot 10^3 < \text{Re}_a < 2 \cdot 10^4$ , Re<sub>w</sub> = Re<sub>a</sub>  $v_a / v_w$ .

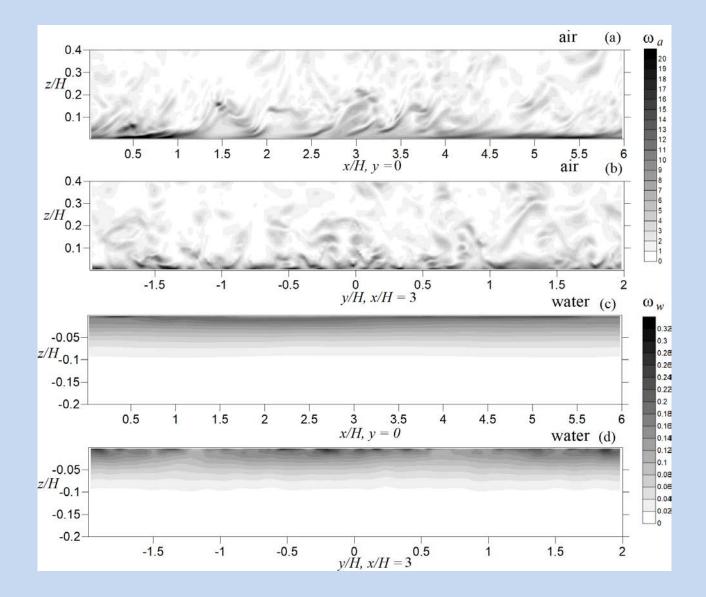
Boundary conditions at the water surface (z = 0):

$$\frac{\rho_a}{\rho_w} \frac{\partial U_{x,y}^a}{\partial z} = \frac{\operatorname{Re}_a}{\operatorname{Re}_w} \frac{\partial U_{x,y}^w}{\partial z}, \quad U_{x,y}^a = U_{x,y}^w, \quad U_z^a = U_z^w = 0, \quad (4)$$

where  $\rho_a / \rho_w \approx 10^{-3}$  - the air/water density ratio.

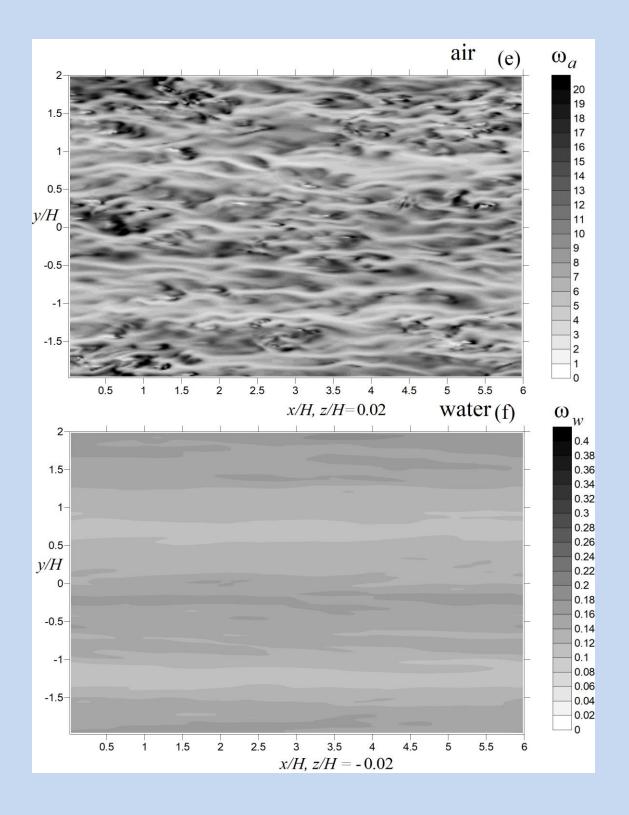
Initially the water is at rest, the air is slightly ( $\sim$ 5%) perturbed.

Vorticity modulus  $\omega_{a,w}$  in central planes (x,z) and (y,z) in air (a,b) and water (c,d) at  $tU_0/H = 400$ . Re<sub>a</sub> =  $1.5 \cdot 10^4$ 

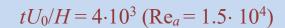


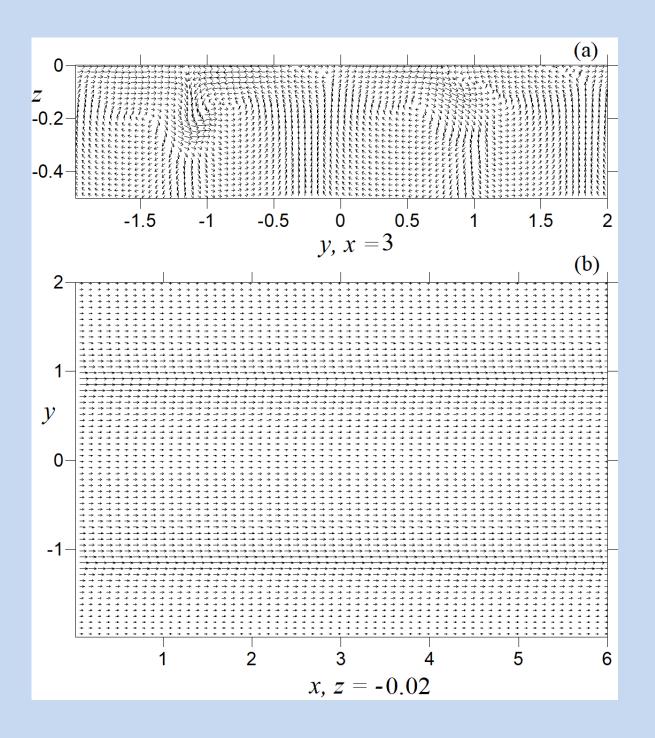
The air-flow is fully-developed and turbulent. The water-flow is quasilaminar,  $\omega_a \gg \omega_w$ , and consists of streamwise-oriented streaks.

Vorticity modulus in horizontal (*x*,*y*) planes in air (e) and water (f) close to the water surface  $(z/H = \pm 0.02)$ 

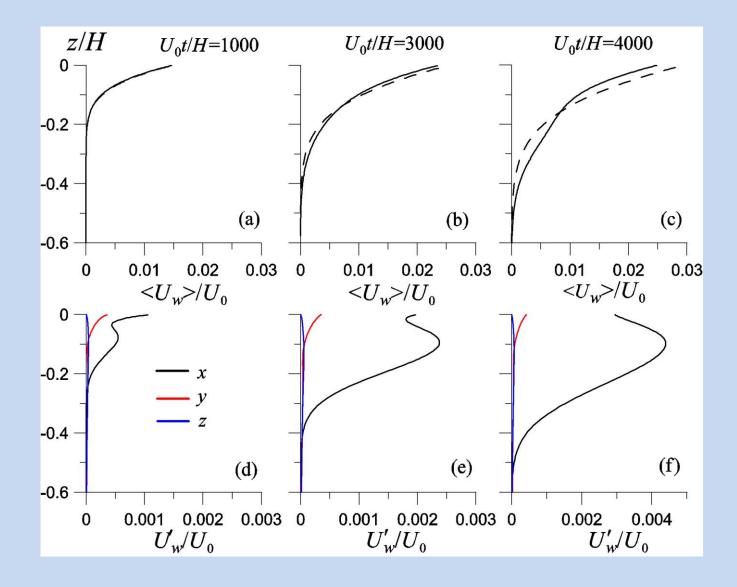


#### Velocity field in water in vertical (y,z) and horizontal (x,y) planes at времени





Vertical profiles of mean drift flow (a,b,c) and fluctuations (d,e,f) at different times. Re<sub>a</sub> =  $1.5 \cdot 10^4$ 

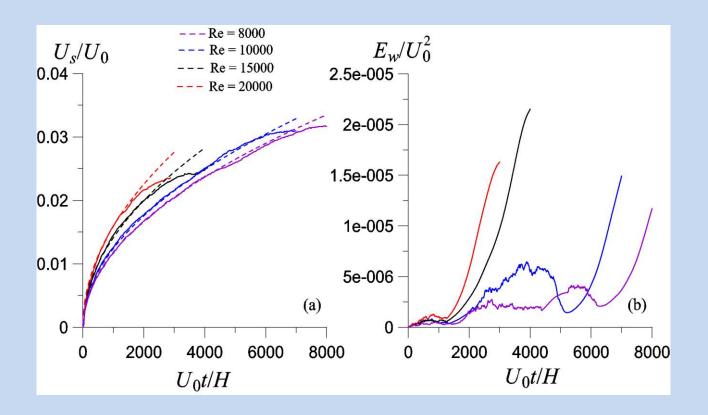


Dashed line is the analytical solution of the 1D Navier-Stokes ("diffusion") equation:

$$U_d^l(z,t) = -\frac{\partial U_d^l(0)}{\partial z} \left\{ 2 \left( \frac{v_w t}{\pi} \right)^{1/2} \exp\left( -\frac{z^2}{4v_w t} \right) + z \operatorname{erfc}\left( -\frac{z}{2\sqrt{v_w t}} \right) \right\}, \quad (5)$$

 $\operatorname{erfc}(z) = \frac{2}{\sqrt{\pi}} \int_{z}^{\infty} \exp(-x^{2}) dx$  $\frac{\partial U_{d}^{l}(0)}{\partial z} = \frac{\rho_{a}}{\rho_{w}} \frac{u_{*}^{2}}{v_{w}}, \text{ where } u_{*} \text{ is the friction velocity of air}$ (6)

Temporal development of surface drift velocity  $U_s$  (a) and kinetic energy of fluctuations in water  $E_w$  (b) for different Re<sub>a</sub>

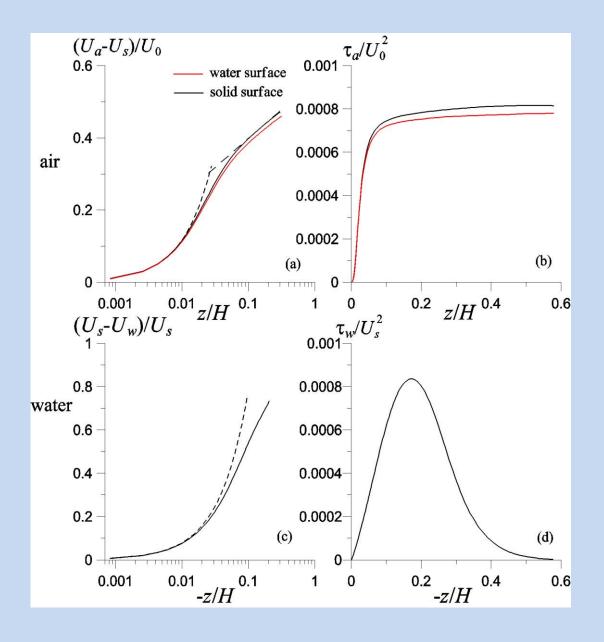


The dashed line is the analytical solution for  $U_s$ :

$$U_{s} = U_{d}^{l}(0,t) = 2u_{*}^{2} \frac{\rho_{a}}{\rho_{w}} \left(\frac{t}{\pi v_{w}}\right)^{1/2}$$
(7)

The analytical solution agrees well with (7) when the fluctuations remain small, until their exponential growth.

Profiles of mean velocity  $U_{a,w}$  and turbulent momentum flux  $\tau_{a,w}$  of air (a,b) and water (c,d). Re<sub>a</sub> =  $1.5 \cdot 10^4$ 



Dotted and dashed lines are for viscous and logarithmic asymptotics

Air: 
$$U_{visc} = \frac{u_*^2 z}{v_a}, \quad U_{log}(z) = 2.5u_* \left( \ln \frac{z u_*}{v_a} + 5 \right).$$

Water:  $U_s^{visc} = U_s + \frac{\partial U_w}{\partial z}\Big|_{z=0} z$ 

## Parameterization of surface drift under low wind

The Reynolds number of the drift flow:

$$\operatorname{Re}_{d} = \frac{U_{s}}{V_{w}} \left( V_{w} t \right)^{1/2} = \frac{2}{\sqrt{\pi}} \frac{u_{*}^{2} t}{V_{w}} \frac{\rho_{a}}{\rho_{w}} \quad . \tag{1}$$

Turbulence in water develops at times  $tU_0/H \approx 1.5 \cdot 10^3$  (Re<sub>a</sub> = 2·10<sup>4</sup>),  $tU_0/H \approx 2.5 \cdot 10^3$ (Re<sub>a</sub> = 1.5·10<sup>4</sup>),  $tU_0/H \approx 4 \cdot 10^3$  (Re<sub>a</sub> = 10<sup>4</sup>),  $tU_0/H \approx 5 \cdot 10^3$  (Re<sub>a</sub> = 8·10<sup>3</sup>)], so for  $\rho_a / \rho_w \approx 10^{-3}$  and  $\nu_a / \nu_w \approx 15$  so that

the critical Reynolds number of the drift flow is defined from DNS:

$$\operatorname{Re}_{d}^{c} \approx 500.$$
 (2)

From (1) the time of transition to turbulence is

$$\frac{t_c}{v_w} \approx \frac{\sqrt{\pi}}{2} \frac{\operatorname{Re}_d^c}{u_*^2} \frac{\rho_w}{\rho_a} , \qquad (3)$$

Then the critical surface drift velocity:

$$U_s^c = \left(\frac{2}{\sqrt{\pi}}\frac{\rho_a}{\rho_w}\operatorname{Re}_d^c\right)^{1/2} u_*, \qquad (4)$$

Then for  $\rho_a / \rho_w \approx 10^{-3}$  and  $\operatorname{Re}_d^c = 500$  (4) gives:

$$U_s^c \approx 0.75 u_*$$

This is an upper estimate of the surface drift attained under low-wind at the stage where the velocity fluctuations are small and the surface remains aerodynamically smooth. For  $U_0 \sim O(1m/s)$  and  $H \sim O(1m)$  the transition time is of order  $t_c \sim O(10^3 s)$ .

#### CONCLUSIONS

The dynamics of a drift flow in the near-surface water layer driven by a turbulent air wind is investigated by direct numerical simulation (DNS). Comparatively low (up to  $2 \cdot 10^4$ ) bulk Reynolds numbers of the air-flow are considered when the air boundary layer is turbulent but velocity fluctuations in the water are sufficiently small and the water surface remains aerodynamically smooth. It is shown that a drift flow develops in the near-surface water layer, and its velocity grows monotonically with time. At long times there develops an instability which leads to saturation of the growth of the drift-velocity. A threshold Reynolds number is defined in DNS under which the drift flow becomes unstable, and a parameterization of the surface drift velocity is formulated in terms of the air-flow friction velocity.

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